

Equilibrium figures

- Historical aspects
 - Collisions + reaccumulation
 - Rubble piles - Agrégats gravitationnels
 - Long spin period, elongated bodies, LASPA
 - Stars, planets and satellites spheroidal
 - Asteroids tri-axial ellipsoids, sometimes very elongated
 - What size barrier for rubble piles?

Gravity

- Different cases in this school
- small g
 - the granular systems on Earth, experiments and natural
- micro μg
 - granular and regolith on surface close to spin barrier v_{esc} , or not so close
- capital G
 - self gravitating bodies
body scale $>$ grain scale

Equilibrium figures

- Observations
- Shape models
 - Farinella
 - Magnusson
 - Holsapple
- Spin barrier
 - Rubble piles
 - Other forces in play ?

Triaxial Equilibrium Ellipsoids among the Asteroids?

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AND

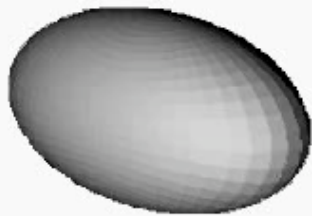
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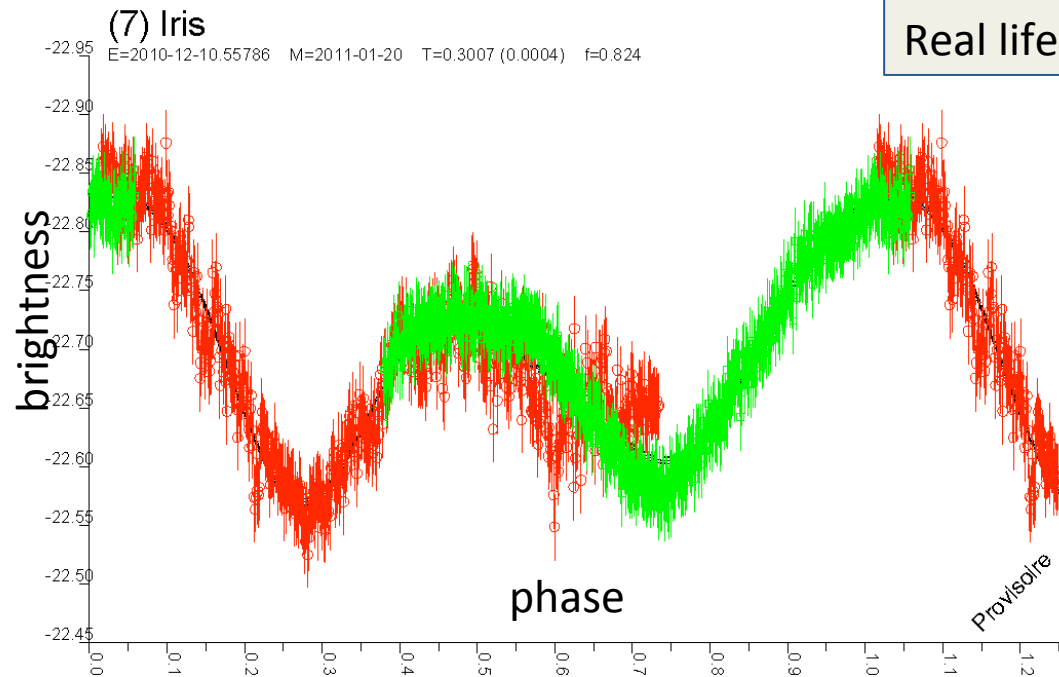
Received October 17, 1980; revised January 27, 1981

Shapes from LC

Theory



Real life

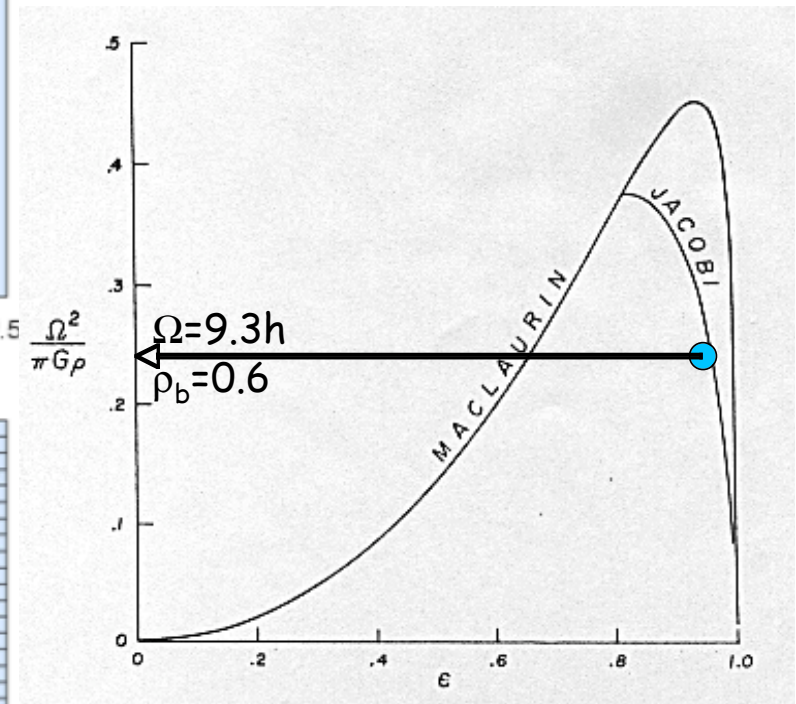
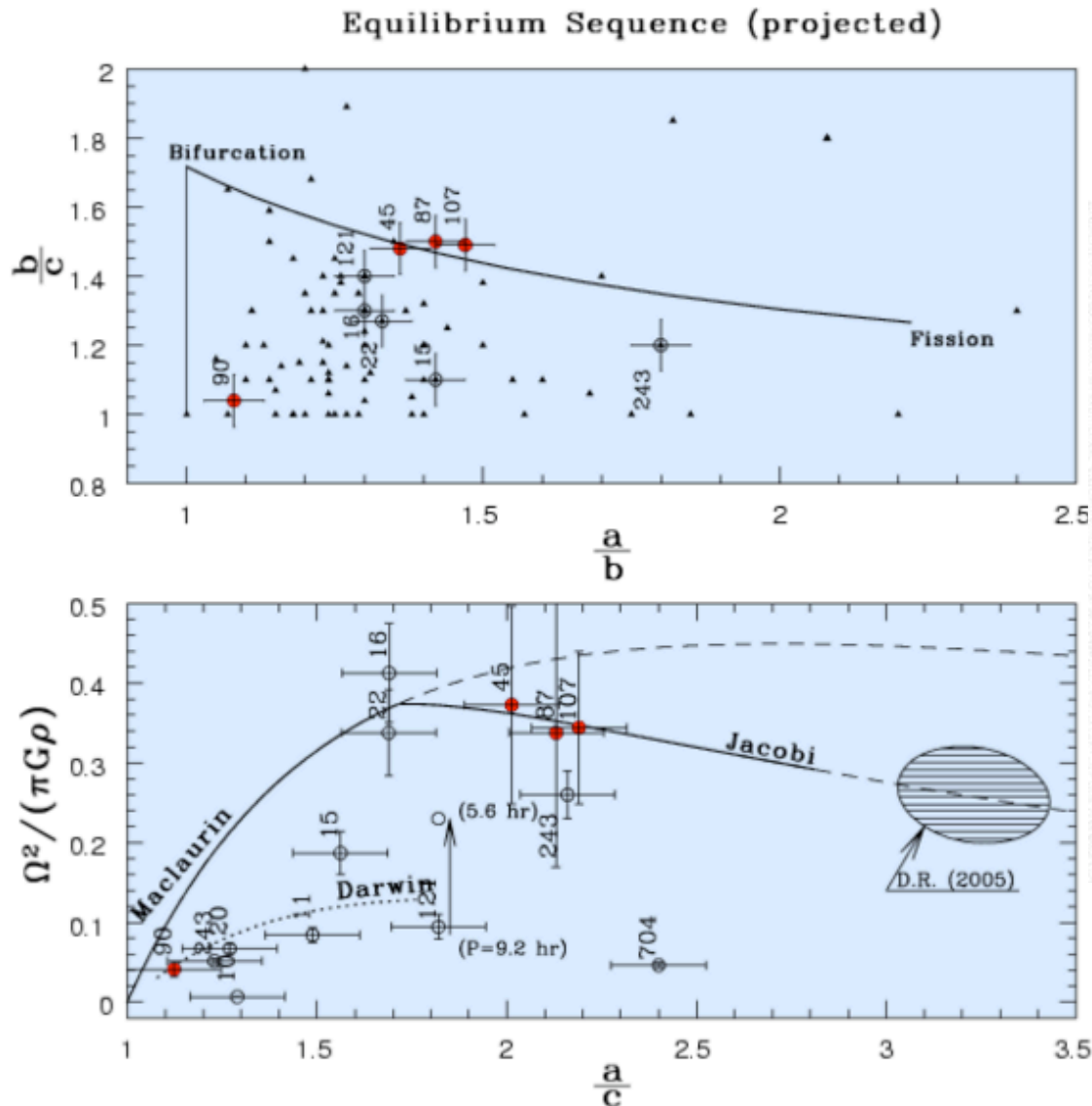


Convex model



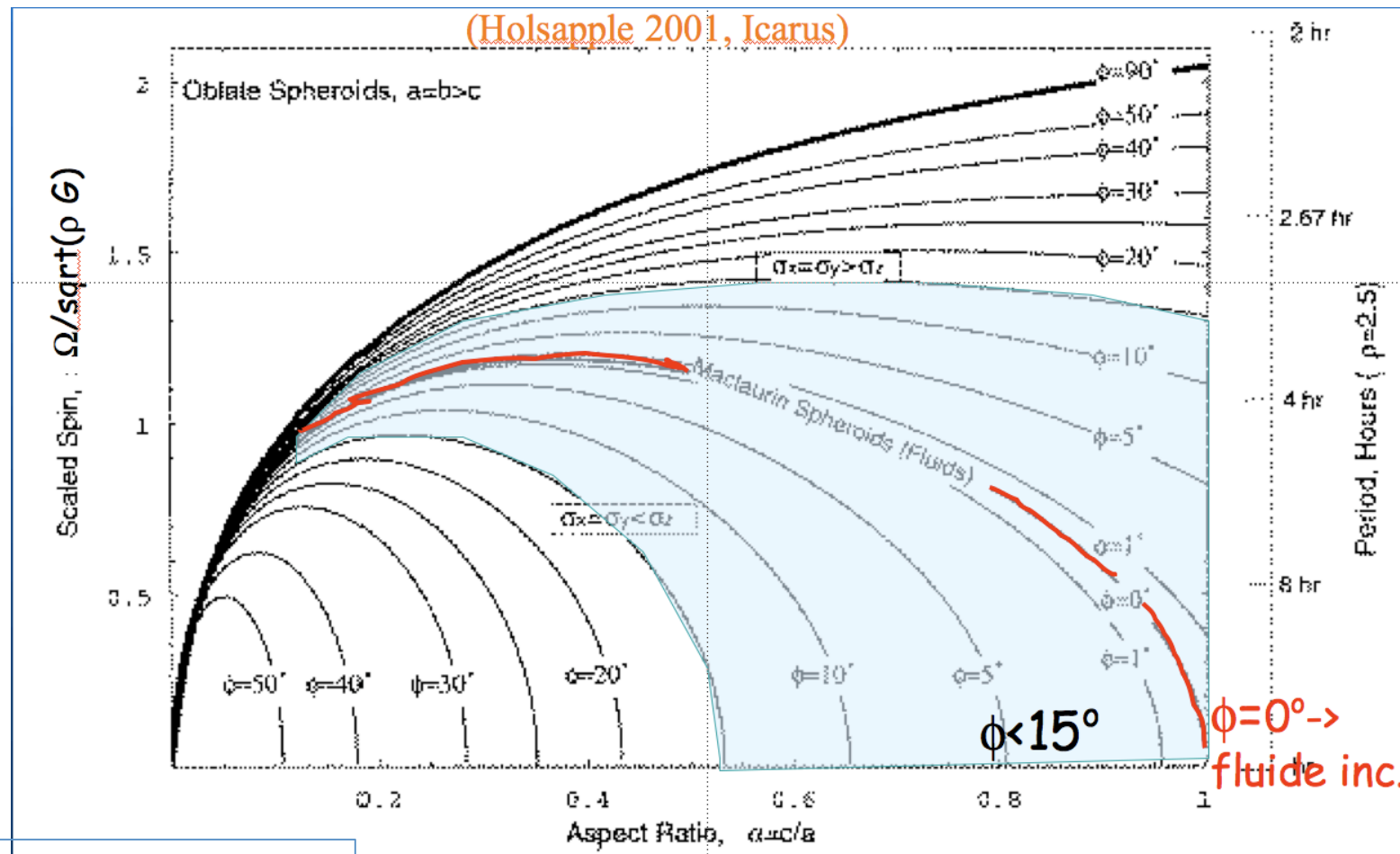
NB: shape, but no size!

- Hydrostatic equilib. (incomp. fluid) not valid



Inferred density of Ausonia

Mohr-Coulomb

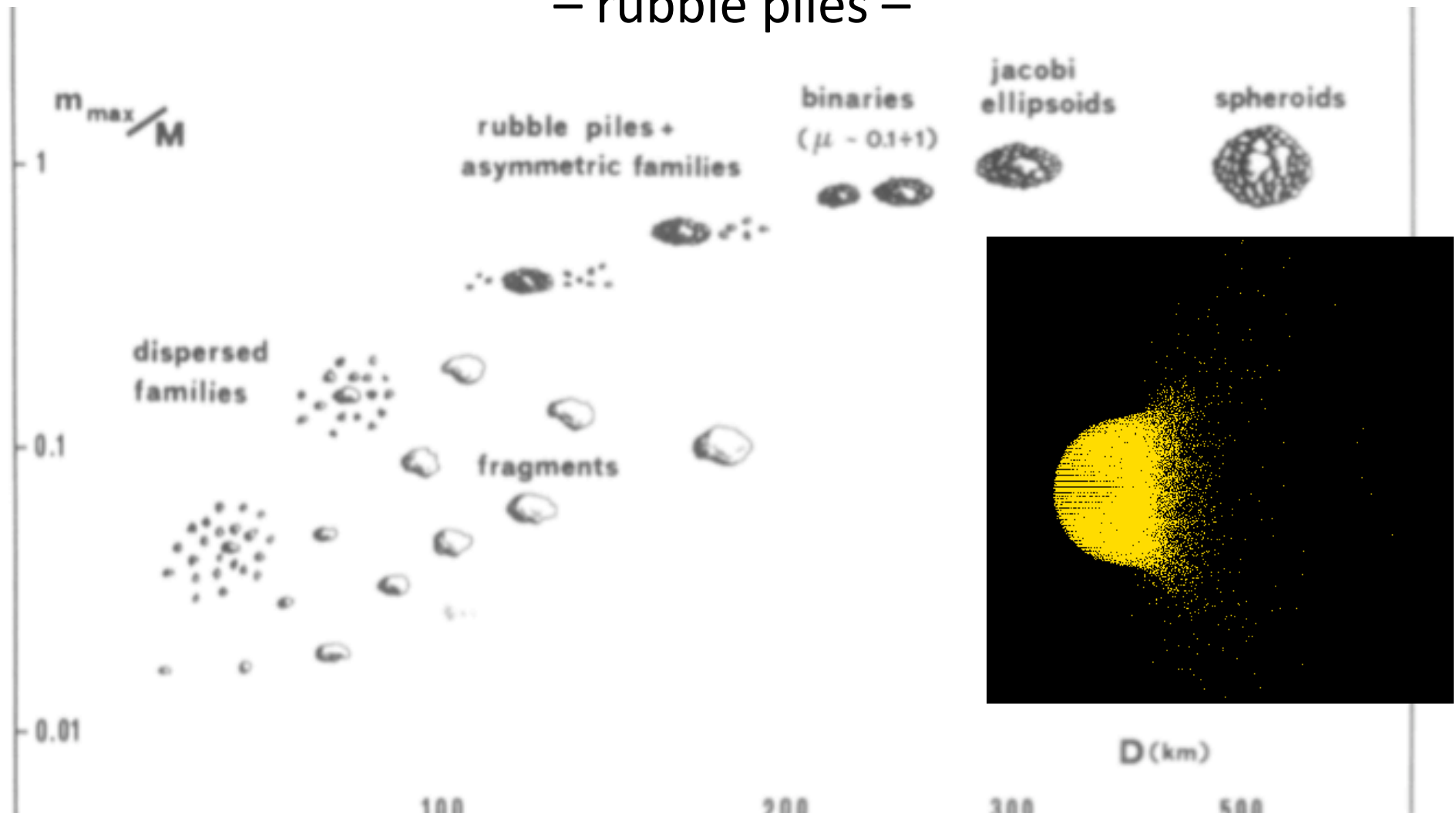


Equilb. +stability

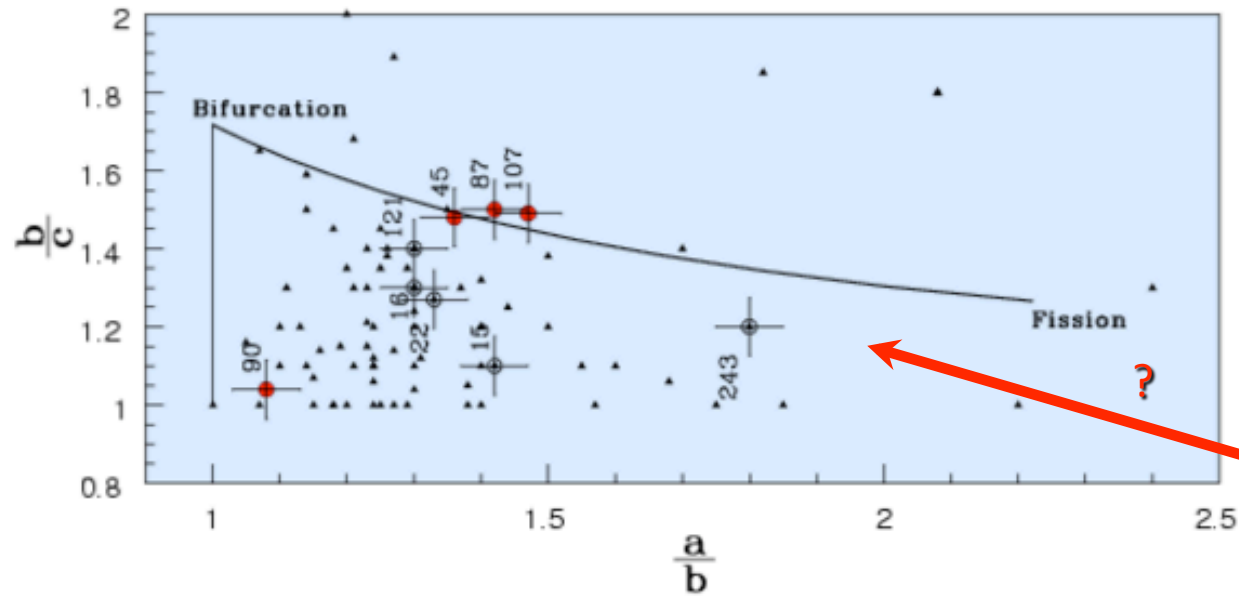
Farinella et al. 1982

Asteroids as outcome of catastrophic collisions

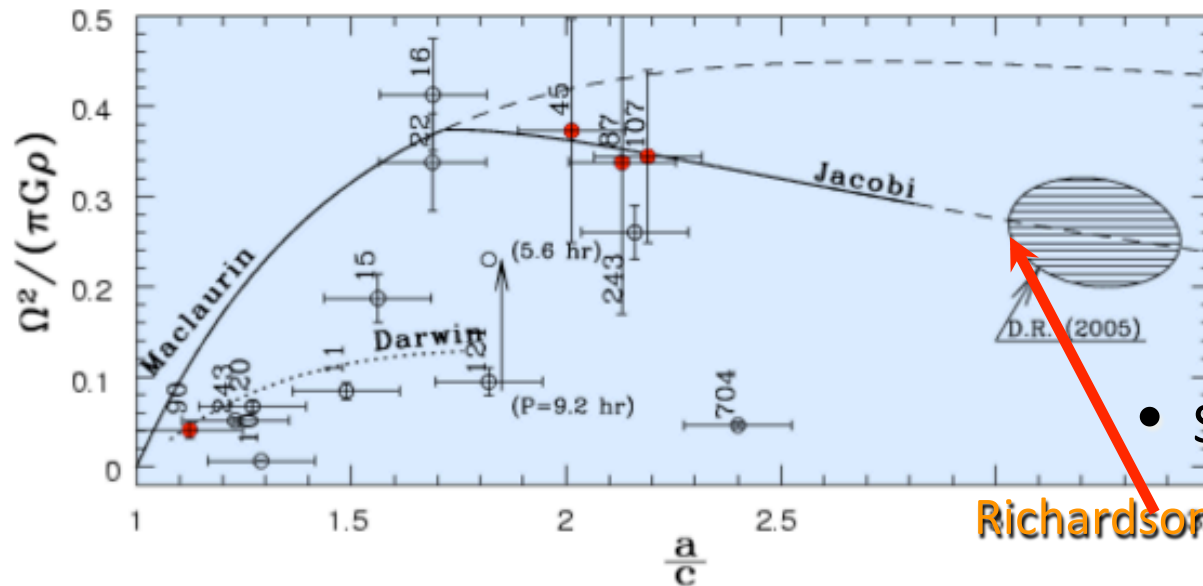
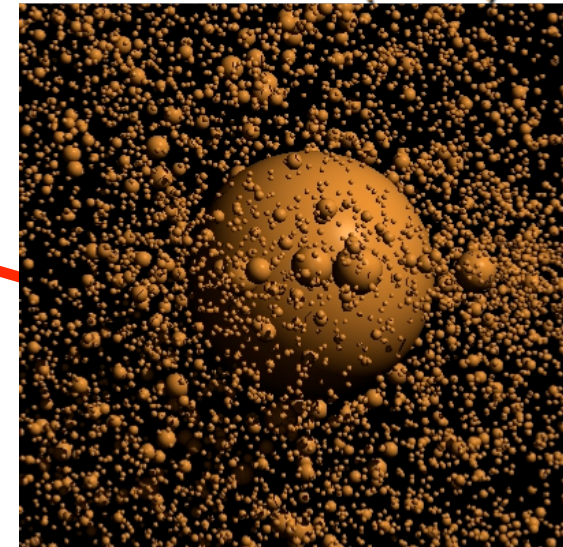
– rubble piles –



- Final result of re-accumulation?



Michel et al. (2001)



- shape not of interest in original simulations

- Spin-up + mas loss

Richardson et al. (2005)

Equilibrium figures

- Binary asteroids
 - Detection, study, not easy
 - MBB, TNB, NEB, Trojans and Centaurs
- Origin - can be diverse
 - Fission (in two)
 - breakup
 - Re-accumulation
 - Others (capture, ...)
 - « The story of the black sheep »

Equilibrium figures

Dobrovolskis (1982) stresses in tri-axial bodies

Shear stress

Tresca

$$|\tau|_{\max} = (\sigma_1 - \sigma_3)/2 > S_0$$

Friction

$$|\tau| = S_0 - \tan(\phi) \sigma$$

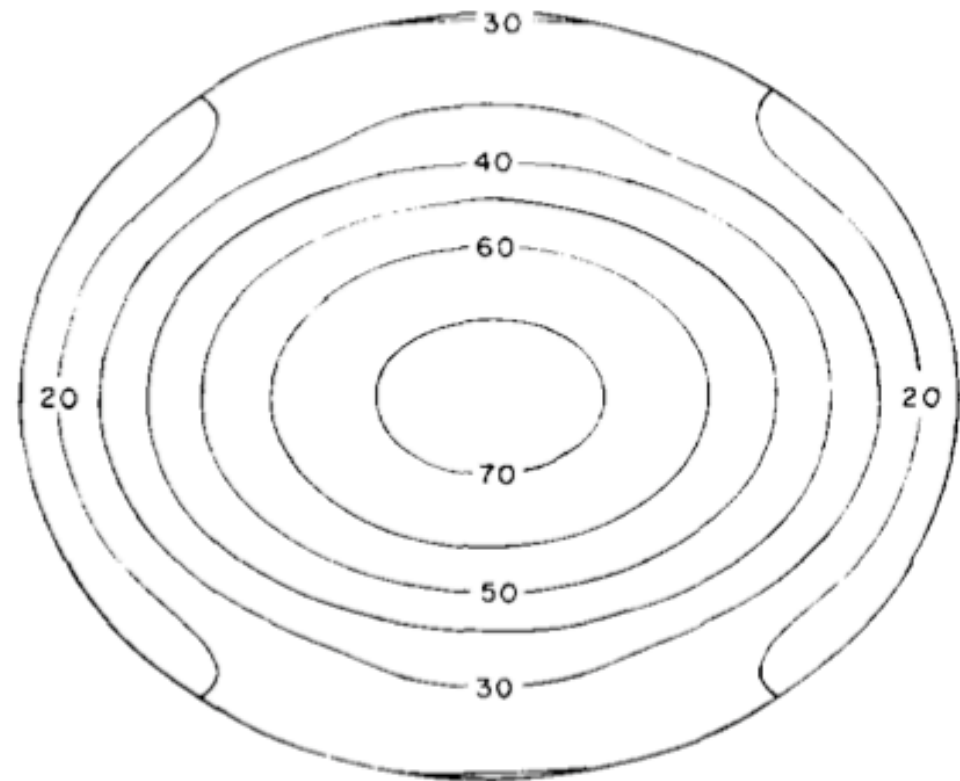


FIG. 4. Contour map of the maximum shear stress $|\tau|_{\max}$ (in millibars) over an equatorial section of Phobos.

Canup (2005)

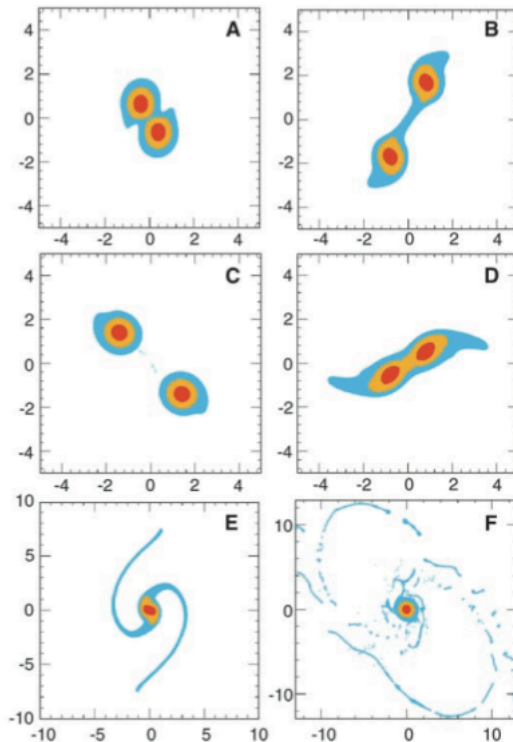
giant impact + reaccumulation

SEARCH ARTICLES

A Giant Impact Origin of Pluto-Charon

Robin M. Canup

Fig. 1. Time series of a potential Pluto-Charon-forming impact yielding a planet-disk system (run 70 in table S1 with $N = 120,000$ particles). Results are shown looking down onto the plane of the impact at times $t = 1.3, 3.2, 7.5, 11.8, 14.5$, and 24.6 hours; units shown are distance in 10^3 km. Color indicates material type (blue, water ice; orange, dunite; red, iron), with all of the particles in the 3D simulation overplotted in order of increasing density. The impacting objects are identical—both are predifferentiated into 40% ice mantles and 60% rock cores by mass with initial surface temperatures set to 150 K, increasing with depth (7) to a central temperature ≈ 800 K. After an initially oblique impact in the counter-clockwise sense (A), the two objects separate (B and C) before recolliding. After the second collision, the denser cores migrate toward the center, as a bar-type mode (36)



forms in the rapidly rotating merged objects (D). From each end of the bar emanate spiral structures (D and E), whose self-gravity acts to transport angular momentum from inner to outer

an ice mantle, rock core, and i and (iii) SIM: 50% serpentine : ice in an undifferentiated mixt jects range from uniform to h tiated, with rock mass fractio and 86% and bulk densities be 2.5 g/cm^3 .

I modeled a variety of imp all capable of providing an ar tum within the range for Plu collision of two nonspinning objects delivers a normalized momentum (I_0)

$$J_{\text{col}} \equiv \frac{L_{\text{col}}}{L} = \sqrt{2} f(\gamma) b'$$

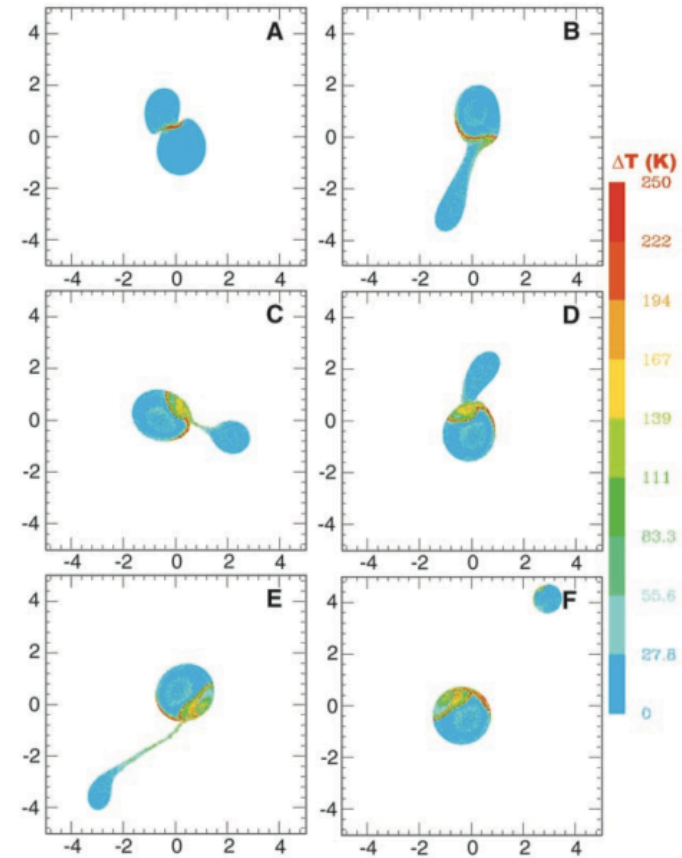
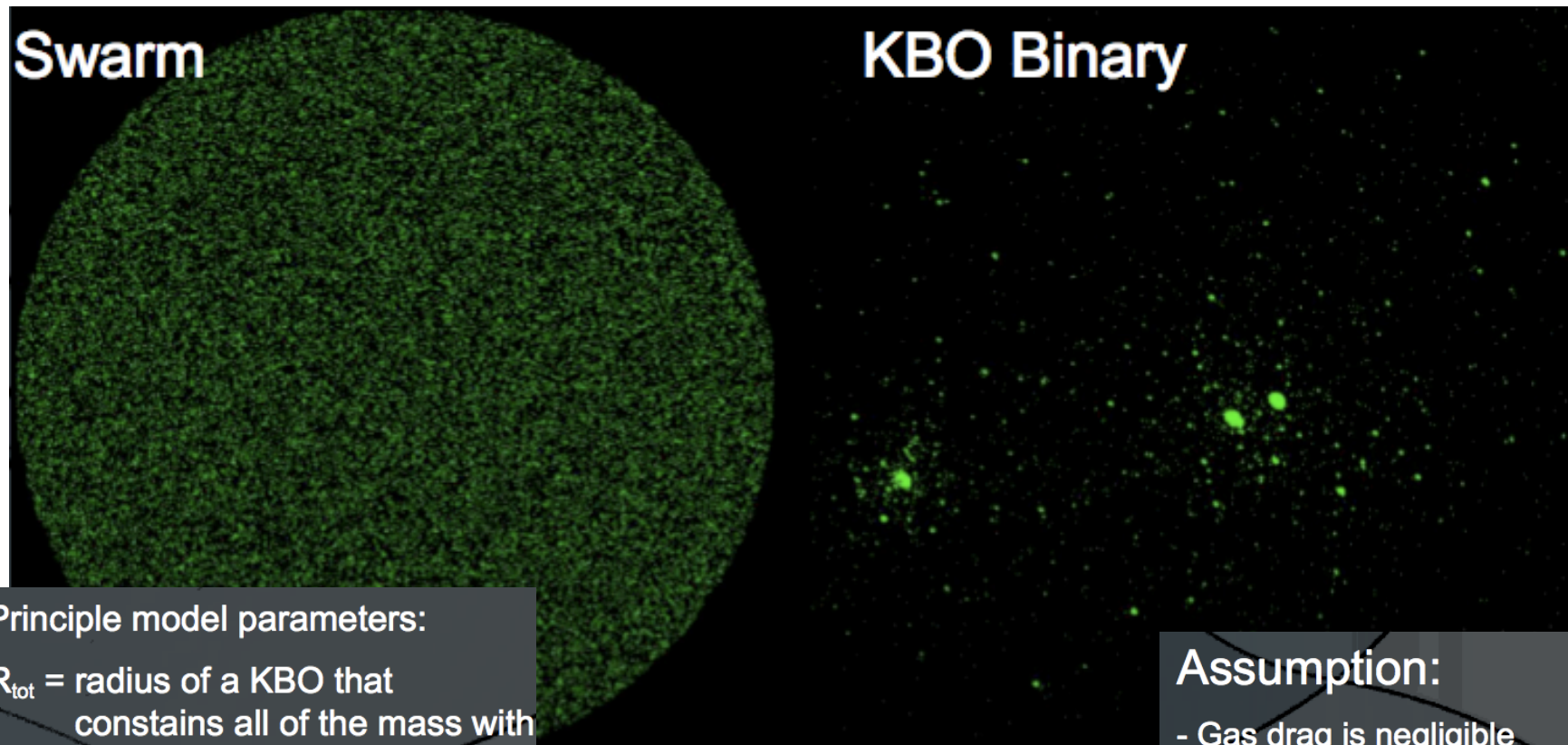


Fig. 2. Time series of a potential Pluto-Charon-forming impact yielding a planet-moon system (run 20 in Table 1 with $N = 20,000$ particles). Results are shown at times $t = 0.9, 3.2, 5.9, 7.5, 11.2$, and 27.5 hours; distances are shown in units of 10^3 km and color scales with the change in temperature in kelvin. The impacting objects have uniform serpentine compositions. After an initially very oblique impact with a 73° impact angle (A), the two objects separate (B and C) and during this period the smaller impactor receives a net torque from the distorted figure of the target. After a second, even more grazing encounter (D), an additional portion of the impactor is accreted onto the planet, while the rest self-contracts into an intact moon containing 12% of the central planet's mass that is again torqued by the ellipsoidal figure of the target (D and E) onto a stable orbit with a semimajor axis of $6.5 R_p$ and an eccentricity of $e = 0.5$. The final moon in (F) is described by 2232 SPH particles.

Nesvorny et al. (2008-2010)

gravitational collapse



Principle model parameters:

R_{tot} = radius of a KBO that contains all of the mass with a density of 1.0 g/cm^3

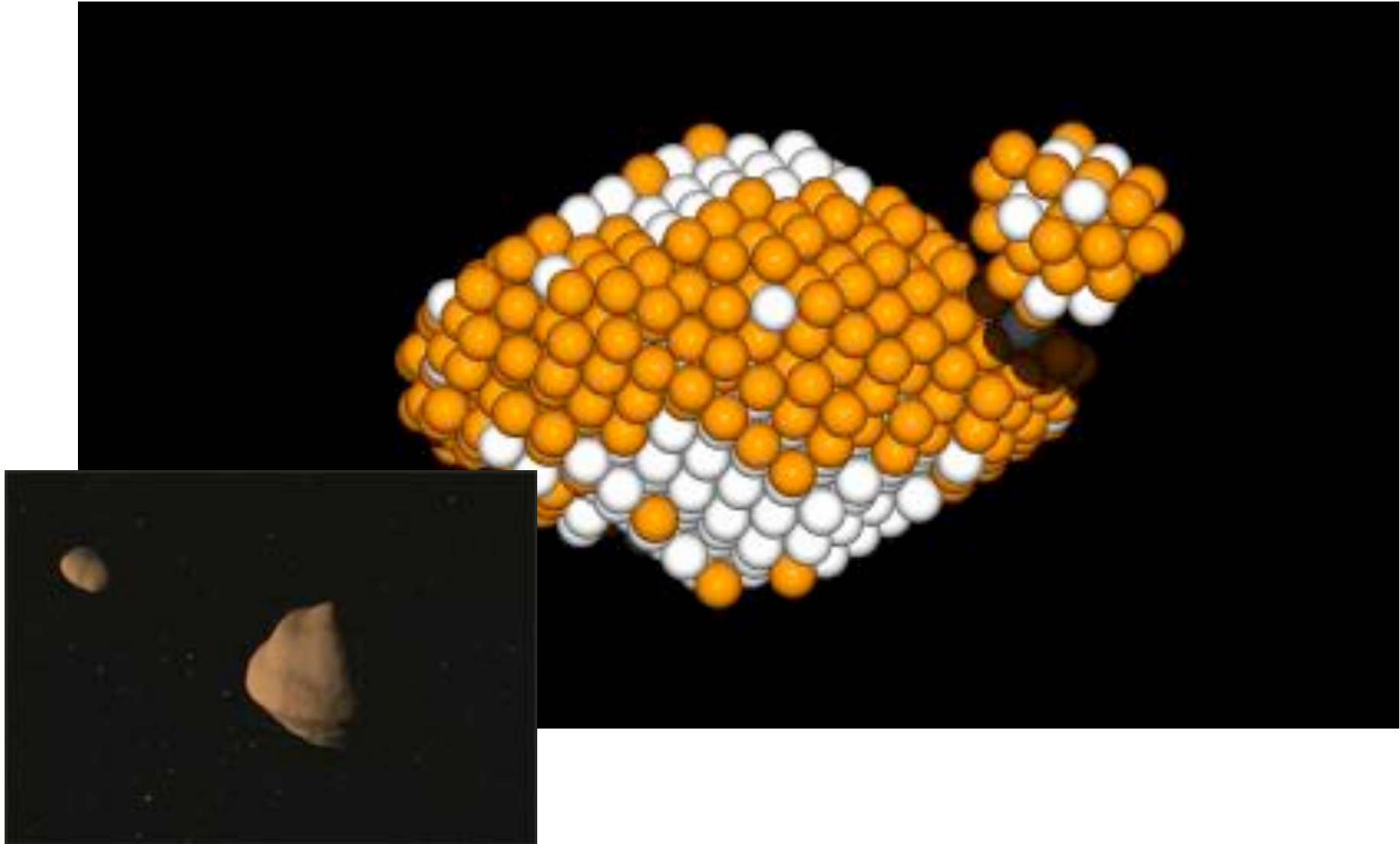
Ω = orbital frequency of swarm around its own center of mass

Assumption:

- Gas drag is negligible
- Solar tides are negligible
- Particle sizes ~ 1 meter

Walsh et al. (2008)

Spin-up



Equilibrium figures

- Binaries
 - Mass + volume => density
 - (hydr. equil.) spin+shape => density
 - =>Test
 - Spin+shape + J2
 - => interior
 - Evolution, tides ?
 - Porosity is uncertain (meteorite analogue)

Shapes & Case studies

- What interior ? what behavior to solicitations (forces, torques, stresses, heat,...)
- What influence of history (tides, collisions, ...)
- Ranges of size different model ?
 - 100m — 10km — 100km

$$\begin{aligned} P(r) &= \frac{2\pi}{3} \rho^2 G R^2 \left(1 - \frac{r^2}{R^2} \right) \\ &= 1.4 \text{ MPa} \left(\frac{\rho}{10^3 \text{ kg/m}^3} \right)^2 \left(\frac{R}{100 \text{ km}} \right)^2 \times \left(1 - \frac{r^2}{R^2} \right) \end{aligned}$$

- 0.1 Pa to 100 Mpa, pressure at centre
- visco-plastic and/or visco-elastic ?

Conclusion

1. bad weather on Wednesday

